FOCAL MECHANISMS OF SOME CRUSTAL EARTHQUAKES THAT OCCURRED IN THE PANNONIAN DEPRESSION (ARAD – SOUTH TIMIŞOARA AREA), THE MOESIAN PLATFORM AND NORTH-DOBROGEAN OROGEN

ZINA MALIȚA, FLORIN RĂDULESCU

National Institute of Research and Development for Earth Physics, P.O. Box MG-2, 077125, Bucharest

The paper presents the results of processing made upon the observation data concerning the P-waves generated by the crustal earthquakes (and a few subcrustal ones) that occurred in the Pannonian Depression (Arad – South Timişoara area), the Moesian Platform (in the Wallachian and Dobrogean sector) and through the North-Dobrogean orogen. Short considerations about the seismicity of these active areas were made, as well as about the important seismotectonic elements.

Key words: seismic shock, focal mechanism, crustal earthquake, tectonic stress, active fault, Pannonian Depression, Moesian Platform, North-Dobrogean Orogen.

1. INTRODUCTION

Many papers (Romanian and foreign) analyzed the problem of earthquake mechanisms located in Romania, particularly those placed in the active region of Vrancea (Enescu, 1962; Constantinescu, Enescu, 1963; Müller *et al.*, 1978; Radu, 1965, 1979; Iosif, Iosif, 1979; Enescu, 1980; Constantinescu, Enescu, 1985; Oncescu, 1987; Crişan, Oncescu, 1987, 1988; Gerner, 1995; Radulian *et al.*, 1996; Ardeleanu *et al.*, 1996; Utale *et al.*, 1994, 1995, 1996; Popescu *et al.*, 1997, 1998).

Some of them were related to earthquakes occurred in other active areas such as: Râmnicu Sărat, Tulcea, Banat, Carpathian Foredeep, especially in the bending zone of the Eastern Carpathians, etc. A recent paper (Radulian *et al.*, 2002) is the compilation of results for 526 earthquake mechanism studies produced in Romania during 1929–1997. While preparing the paper, 26 sources of information were used, consisting in published papers (Romanian and foreign) and scientific reports from the archives of the National Institute of Research and Development for Earth Physics. Information is placed in the Vrancea region (subcrustal and crustal area), Banat, Danubian area (western the Southern Carpathians), the Făgăraş–Câmpulung region, Bârlad and Predobrogean Depression. The fault plane solutions were determined, based on P wave polarity (the first arrivals) and in some cases by waveform inversion of local earthquakes and teleshocks. All solutions were determined before 1980 and have been recalculated by Aki and Richards conventions (1980).

Given the diversity of the experts who compiled the survey, the authors of this paper proposed the resumption of this major problem with seismotectonic implications, based on a uniform treatment of primary observation data using the same processing method. Before the presentation of the mechanism parameters, brief consideration of the seismicity and seismotectonics of investigated areas were presented.

2. PANNONIAN DEPRESSION (ARAD – SOUTH TIMIŞOARA)

The area included in our analysis is limited to the following geographic coordinates: latitude $-45^{\circ}-46^{\circ}N$ and longitude $-20^{\circ}-22^{\circ}E$. The first mentions on seismicity of the region were made by Mathei M. Drăghiceanu in his study published in 1896, *Monographie des tremblements de terre*

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de Roumanie et des pays environnants, referring to earthquakes produced during the 1892–1894 period in eastern and western part of the country (Banat, Drobeta-Turnu Severin).

Later, Prof. Ion Atanasiu (1949) appointed the seismic activity from this sector of the Pannonian Basin in the Banatic category, with epicenters at Banloc, Vinga, Sânnicolau Mare, Şag, Timişoara and Periam. Subsequent observations revealed large epicentral intensity (VII and VIII degrees on the Mercalli scale) in Banloc, Liebling–Voiteg, Şag–Parţa, Timişoara, Jimbolia, Periam, Arad and Sânnicolau Mare (Oros, 1991). The author mentions the following strong earthquakes: 1915 (Banloc), 1959 (Şag– Parţa), 1979 (Sânmihaiu Român), 1988 (Buziaş) and 1991 (Banloc and Voiteg). In 240 years (1766–2006), 269 earthquakes were located in this region, 6 of them with $M \ge 5.0$ (1879, 1901, 1959 and three shocks in 1991), 51 earthquakes with M = 4.0 to 4.9 and 92 with M = 3.0-3.9.

The catalogue edited for the region indicated several periods of time with low seismic activity, namely: 1916–1920, 1928–1932, 1937–1940, 1942–1952, 1961–1967, 1969–1972, 1975–1977. In the last period seismicity analysis revealed a number of active periods in terms of the seismic activity, namely: July–December 1991, May–August 1995, October 1999, July–August 2000 and November 2002. The distribution of epicenters of the investigated area is shown in Figs. 1 and 2, showing the 2D and 3D image representation of the seismic foci.



Fig. 1 – Earthquake epicenters distribution that occurred during the 984–2006 period (tectonic background by Radu, Oros, 1991).



Fig. 2 – Earthquake epicenters distribution, isobath map and 3-D image of the seismic foci surface of Banat.

Referring to the most active sequence of 1991, within 6 months there have been produced 108 seismic shocks (3 with M > 5.0 and 7 with $M \ge 4.0$). The three major earthquakes have epicenters in Banloc and Voiteg areas (South of Timişoara).

Geological and geophysical (gravity) studies evidenced several major fractures of the basement, E–W and NW–SE oriented, and the existence of vertical and horizontal movements along these (Visarion *et al.*, 1979). These faults are major tectonic blocks bounded with raising and sinking positions. In a seismotectonic study, Polonic and Maliţa (1997) noted that seismic activity is concentrated along Timişoara–Nădlac and Buziaş–Arad faults, that border the Caransebeş and Sânnicolau Mare grabens.

Among these, the elevation Battonya–Buziaş block is individualized; the Lucareț and N Timișoara faults are mentioned, the first separating (after some authors) the Pannonian Block (North) from the Geto-Danubian block (south of Lucareț fault).

The Timişoara area is characterized by large positive values (80–90 mWm⁻²) of geothermal anomalies (Demetrescu *et al.*, 1991). The geothermal studies reveal that the heat flow source observed at the surface is placed in the upper mantle. After these authors, the temperature distribution in lithosphere of the region is a major factor determining the increase of thermoelastic stress, thus, the generation of normal earthquakes (crustal).

The stress indicators, the focal mechanisms of the earthquakes and *in situ* measurements (in wells) have shown that the accumulated stress in this active area is the compressive type, being caused by the convergence movement and rotation of the Adriatic microplate. According to Polonic and Maliţa (1997), the crustal earthquakes of this region are caused by the compressive forces acting between the Pannonian, Geto-Danubian and Banloc blocks, along the Lucareţ reverse fault.

In Table 1, the focal mechanism solutions for 60 crustal earthquakes produced in the period 1991–2001 are shown. The processing observation data was made with the program Wickens and Hodgson (1967), improved by Oncescu (1980).

In determining the location parameters of these earthquakes (epicenter coordinates, hypocenter depth) information from the archives of National Institute of Research and Development for Earth Physics and the data provided by the National Center in Denver, Colorado, USA (National Earthquake Information Center) and the International Seismological Centre Newbary, Centre, UK (International Seismological Edinburgh, Newbary) have been used. In connection with the observation data, in order to obtain good results, the distribution of azimuth around the epicenter, in at least seven observatories, seismological where focal mechanism solutions were primary calculated. Mechanisms of events with less than seven stations were estimated based on their polarities consistency mechanisms other events that could be determined with a high degree of confidence.

Seismic foci depths were 5-35 km in the middle crust (in lower domain of the granitic layer), here the Conrad discontinuity is considered to be at depths of 14-15 km.

Focal mechanism solutions indicate the faulting plane with two-average directions: the first, N41^oW (for 32 shocks) and the second, N47^oE (for 28 shocks); average inclination of the breaking plane was 56^o. The second (Plane 2) showed also mean directions: one, N50^oE (for 31 shocks) and the other, N48^oW (for 29 shocks) and average inclination – 60° . Faulting type was 43 reverse faults and 17 faults strike-slip.

Concerning the tectonic stress, data processing indicated following compressive forces (P) and tension (T) orientations: for P axes, two prevailing directions around N37⁰W (for 31 shocks) and N45⁰E (for 29 shocks) and for T axes – N44⁰W (for 35 shocks) and N52⁰E (for 25 shocks). The predominant inclination of the compressive force P was around 17^{0} and the tension forces (T), around 47^{0} (Fig. 3).

Strong earthquakes on the Banloc and Voiteg of July 12, 1991 (Mw = 5.6) and December 2, 1991 (Mw = 5.5) were the fault plane azimuth WNW–SEE; compressive forces operated NW– SE, and the tension on NE–SW direction. Seismic foci depths were 11 km and 9 km and both faults were the strike-slip fault type.

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	Dip	27	22	23	39	33	42	46	37	63	54	44	24	36	68	46	40	09	82	26	46	42	42	37	77	29	43	60	27	46	52
8	Az	26	249	181	9	231	38	200	185	146	139	174	197	165	89	232	156	129	294	77	174	219	98	70	83	16	164	304	217	159	203
	Dip	25	64	59	43	49	38	34	51	27	34	44	59	45	22	34	44	22	7	64	22	44	24	33	10	57	19	26	53	13	27
Ъ	Az	24	131	308	118	349	287	307	289	54	42	72	321	50	359	341	47	31	204	339	281	113	209	187	351	128	56	207	92	261	306
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2	Rake	21	155	31	165	171	129	55	36	53	39	26	176	136	60	160	135	46	66	16	166	22	66	57	74	32	109	46	143	168	62
PLANE	Dip	20	84	88	45	50	80	77	73	48	46	50	59	82	48	36	77	32	49	69	23	48	88	89	37	80	89	39	88	13	75
Ч	Strike	19	82	178	255	127	226	191	163	119	98	121	66	353	68	118	349	06	120	26	69	160	97	69	70	360	344	267	40	57	195
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	MG	11															3,7								3,4	3,5	3,2	3,2	3,4	3,4	3,3
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	Long.E	7	21,06	21,05	20,99	21,17	21,08	21,13	21,07	21,12	21,12	21,05	21,10	21,24	21,10	21,13	21,06		21,06	21,13	21,15	21,15	21,28	20,97	20,92	20,98	21,00	21,28	-	20,94	21,18
	Lat.N L	6	45,87	45,36	45,59	45,45	45,50	45,35	45,40	45,43	45,34	45,31	45,32	-	45,34	45,33	45,44	45,26	45,40	45,43	45,46		45,48	45,45	45,31	45,50	45,28	45,35	45,51	45,42	45,26
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Origin	Time hh:mm:ss	5	15:27:32	10:42:21	13:55:15	16:29:08	05:10:34	17:56:57	17:03:26	23:59:31	01:19:52	01:27:32	05:24:20	03:36:29	03:58:47	11:22:53	12:57:24	00:51:04	15:04:46	21:25:49	04:59:39	23:36:00	01:34:33	23:01:44	07:58:03	07:28:50	22:10:11	12:12:46	07:45:18	07:34:05	11:30:07
	Day	4	26	12	12	12	13	13	14	14	19	19	19	20	20	31	1	2	9	11	12	14	15	18	29	7	11	13	18	19	25
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The focal mechanisms solutions of some earthquakes occured in the Pannonian Basin (Arad - South Timişoara area)

Table 1

88	154	59	83	33	140	47	177	124	340	94	336	328	100	32	236	123	287	167	165	168	37	353	135	139	234	178	111	280	20	272	303	•
27	61	44	38	85	82	42	43	50	27	52	61	18	49	73	65	73	61	60	35	74	24	52	31	64	64	60	49	46	25	00	18	
26	35	208	300	216	50	176	282	300	210	234	156	126	234	196	65	244	121	331	350	292	230	181	328	47	74	292	327	175	222	132	169	
55	15	42	46	ъ	0	35	16	37	52	31	29	70	31	16	24	6	28	29	55	6	99	37	58	~	25	13	35	14	63	80	65	
24	298	313	189	126	320	287	27	38	83	336	246	236	339	287	334	337	28	65	256	24	130	88	228	316	340	29	223	73	114	3	39	
33	24	16	19	0	∞	28	43	11	25	20	0	7	24	4	3	14	9	2	2	13	S	4	9	26	∞	27	19	40	9	9	17	
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21	106	46	16	83	6	55	74	146	142	148	52	18	58	114	121	107	49	55	152	79	14	135	19	87	52	128	127	104	12	1	155	•
8	71	73	48	45	53	82	6	47	89	37	52	82	76	43	53	32	46	58	68	59	70	59	64	19	43	22	72	87	66	80	89	
19	220	192	233	211	230	170	282	165	31	106	130	103	225	34	265	79	90	311	205	286	176	28	276	44	44	142	160	356	160	48	351	
18	i.s.	i.d.	i.d.		-	d.i.d	d.i.d	i.s.	d.i.s.	i.s.	i.d.	d.i.d	i.d.	i.s.	i.s.	ii.	.b.i	i.d.	i.s.	i.	i.d.	i.s.	i.d.	ľ.	i.d.		i.s.	d.i.s.	i.d.	d.i.d	d.i.s.	
17	50	156	137	97	89	166	179	49	-	58	128	172	155	69	55	80	125	132	25	108	159	40	153	91	120	76	28	12	156	170	~	
16	25	46	78	45	37	36	16	99	52	72	52	72	2	52	47	00	57	46	29	33	17	53	73	71	57	73	41	14	79	89	65	,
15	357	84	132	41	49	69	191	280	122	223	1	10	114	183	39	239	322	183	306	127	81	146	177	228	271	289	272	67	66	317	81	
14							3,2				4,5				4,8	4,8	4,1	-				3,8							3			
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9										3,9		3,6	4,6		4,6					4,4					4,8	3,2						
σ						5'2	3,4				3,8							_								_	2,8	3,3	3,2	3,1	3,3	
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7	21,10	21,16	21,25	21,17	21,01	21,11	21,06	21,78	21,09	21,65	21,56	21,32	21,63	21,50	20,95	21,03	20,95	21,21	21,38	21,19	21,00	21,01	21,29	21,00	21,02	20,84	21,24	21,03	21,24	21,33	20,96	
9	45,23	45,38	45,35	45,49	45,36	45,49	45,44	45,56	45,55	45,92	45,90	45,83	45,95	45,90	45,56	45,55	45,53	45,55	45,45	45,55	45,49	45,47	45,94	45,54	45,62	45,86	45,59	45,58	45,62	45,69	45,33	,
5	14:37:51	15:02:14	16:38:42	02:16:31	23:36:36	08:49:40	10:52:35	15:04:00	12:02:04	10:24:12	03:12:22	11:43:06	20:33:21	15:46:44	09:34:06	21:05:14	08:23:05	08:57:57	23:20:30	13:49:28	23:31:27	01:42:51	18:50:35	19:34:39	09:13:28	04:43:40	07:23:36	22:14:01	23:04:11	15:44:48	21:50:45	
			16:2	02:1		08:4	10:5	15:0		10:2		11:4	20:5				08:2	08:5	23:2				18:5	19:5	09:1	04:4	_		23:C	15:4	21:5	
4	17	17	24	21	8	2	2	2	13	18	19	21	2	24	19	33	14	8	4	13	13	15	12	12	24	<i>ლ</i>	22	16	5	20	2	
m	10	10	10	11	11	12	12	12	12	12	12	12	3	5	12	12	1	5	9	10	10	10	11	11	<i>с</i> о	10	2	∞	10	9	00	
2	1991	1991	1991	1991	1991	1991	1991	1991	1991	1991	1991	1991	1992	1992	1992	1992	1993	1993	1993	1994	1994	1994	1994	1994	1996	1999	2000	2000	2000	2001	2001	,
-	30	31	32	33	2	35	36	37	38	39	40	41	42	43	44	45	46	47	48	49	50	51	52	53	\$	55	56	57	58	59	60	,

 Table I (continued)

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 27
 28

oblique fault; **i.d.** – reverse right-lateral oblique fault; **i.s.** – reverse left-lateral oblique fault; **d.n.d.** – right-lateral strike-slip; **d.n.s.** – left-lateral strike-slip; **d.n.s.** – left-lateral strike-slip; **d.i.s.** – left-lateral strike-slip; **d.i.s.** –



Fig. 3 – The dip (a) and azimuth (b) of P and T axes of earthquakes occurred in Banat.

In the Jebel–Banloc active area, moderate magnitude earthquakes in 1915 (Mw = 4.3 to 4.8), 1936 (Mw = 4.8), 1960 (Mw = 4.2) and 1980 (mb = 4.2) occurred. The active region is located in the sedimentary basin south of Sânnicolau Mare Graben.

3. MOESIAN PLATFORM

The investigated region includes the sector of the Moesian Platform and a part of the Carpathian Foredeep. Seismic data have been extended to 44.8° N latitude, and in south to latitude 43.5° N and between meridians 22.5° and 29° E. The catalogue includes the period during 1276–2007, 629 earthquakes, of which 565 produced in Romania, 63 in northern Bulgaria and Black Sea and a seismic shock in Serbia. In Fig. 4 the distribution of the earthquake epicenters within the area referred is shown; a seismic foci 3D representation is shown in Fig. 5.

Earthquake epicenters distribution in this area reveals that the western sector (Wallachian) is much less seismically active than the eastern sector (Dobrogean) located eastern of the Intramoesian Fault. Temporal analysis of the seismic shocks produced in the XXth century has shown a significant decreasing of seismic activity in some periods: 1917–1922, 1924–1929, 1931–1941, 1943–1955, 1957–1959, 1961–1966, 1968–1974 and 1978–1979. It's worth mentioning an increasing of the seismic activity in other years as: 1977 (18 earthquakes), 1982 (43 earthquakes), 1995 (22 earthquakes), 1996 (21 earthquakes), 1999 (29 earthquakes), 2005 (60 earthquakes) and 2006 (48 earthquakes).





Fig. 5 – Earthquake epicenters distribution, isobath map and 3-D image of the seismic foci surface of the Moesian Platform.

On the Carpathian Foredeep area, it seems that, in time, it was a migration of the seismic activity from east to west, from an earthquake of magnitude 4.5 on July 9, 1912 (the epicenter west of Râmnicu-Vâlcea) to a seismic shock of 5.2 magnitude on June 20, 1943 (north of Târgu-Jiu) and July 18, 1991, earthquake with Ms = 5.6 (north-west of Drobeta-Turnu Severin), the last one located in the Southern Carpathians. Most of the XXth century earthquakes occurred in

the range 0-10 km deep (about 65%), the rest (about 35%) at depths of 11-35 km. Thus, the earthquakes mostly occur in the sedimentary crust (or in the upper crystalline, consolidated crust), few of them have focus in the middle crust and others in the lower part of the consolidated crust.

The following can be distinguished as active seismic areas in western sector of the platform itself as: Alexandria–Chiriacu, Turnu Măgurele– Caracal and Calafat–Strehaia North. In the transition zone to the Carpathian Foredeep area it has been shown some areas with seismic potential as Ludeşti–Cobia, Tărtăşeşti–Butimanu and Târgu-Jiu–Turnu Severin. North-western of Turnu Severin outlines a very active area, located in the Southern Carpathians. Here, on a relatively small area (about 2500 s/qkm), numerous seismic events occurred, of low magnitude, which culminated in the July 18, 1991 earthquake (Ms = 5.6, depth 11.6 km).

Correlation with regional tectonic elements indicates the connection of these earthquakes with Cerna fault. The great depth (33 km) of these earthquakes shows the extension of this major crustal fault in the lower crystalline crust; previous seismic studies indicating a thick crust of about 40 km.

Among the earthquakes produced in the central sector of the platform, the Alexandria area one (on February 7, 2001, M = 3.2) was optimal for the focal mechanism determination. The analysis showed the fault plane oriented N25⁰W (dip 42⁰); depth focus (33 km) places, respectively in the subcrustal domain (Table 2).

The seismic activity of the eastern sector (Dobrogea) of the Moesian Platform is well known, especially after strong earthquakes on 1960 and 1967, with magnitudes (Mw) of 5.4 and 5.0 respectively (Iosif, 1960; Iosif, Iosif, 1971). An increasing seismic activity was observed after the Vrancea earthquake of March 4, 1977, so at the end of 1977, there have been 23 earthquakes in this area, five of them having magnitudes of 3.6 to 3.9 degrees on the Richter scale (Cornea, Polonic, 1979). Macroseismic observations first made by I. Atanasiu (1949) have shown in this part of the platform many epicenters and lines of "seismic sensitivity" interpreted as representing the structural alignments of the crystalline basement.

Among the earthquakes produced in the 1960–2007 period, it was possible to determine the focal mechanism only for 75 (Table 2). From these, we mention the earthquakes of January 4, 1960, from Căzănești (Ms = 5.4) and February 27, 1967 at Rădulești (Ms = 5.0), with foci at 40 km, and 42 km, respectively.

Determinations of 63 earthquakes produced in the Dobrogea (east of 26^{0} E meridiane) showed one or two predominant faulting directions, one – N39.6⁰E (for 32 earthquakes) and another – N49.3⁰W (for 31 earthquakes) and inclinations of 59.4⁰ and 55.6⁰, respectively. For other breaking plane the prevailing direction was N35.7⁰ E (for 39 shocks) and N 44.9⁰ W (for 24 events) and prevailing inclinations of 55.8⁰ and 58.6⁰, respectively.

The types of faulting derived from the focal mechanism solutions were the reverse faults and strike-slip faults.

Concerning the tectonic stress, the axis of compression forces (P) presented two prevailing directions, quasi NW–SE (N49, 5^{0} W) and NE–SW (N 41.6⁰ E).

The dips of the compressive forces were almost horizontal, with average dips of 20° and 18° , compared with the tension axis (T) which operated to dips of 41° . Tension stress occurrs around the direction of N42.5°E and N62.5°W. In Fig. 6 were represented the azimuths and dips of the P and T axes for all 75 earthquakes in the investigated area of the platform.

In conclusion, the fault plane solutions from the Dobrogea sector have shown the existence of two faulting systems (already known from previous geophysical works), one mainly oriented NW–SE and the other side, NE–SW. The first category includes the major active faults (from west to east): Intramoesic fault and Nehoiu– Smeeni–Dragalina, Slobozia–Feteşti and Capidava Ovidiu faults. The group of secondary faults is composed of Bărăitaru–Făurei–Oprişeneşti, Urziceni–Jugureanu, Ileana–Colelia, Cartojani– Grădinari and Videle–Bălăria faults (from north to south).

The events analyzed in the 1960–2007 period have shown the existence of a compressive regime, oriented NW–SE, the same tectonic stress appeared too at the earthquake occurred at Rădulești on February 27, 1967 (Ms = 5.0). Along the Intramoesian fault were located only a few earthquakes. Also, more seismic shocks were related to fracture Nehoiu–Smeeni–Dragalina, this fault seems to be related to the shock from January 4, 1960 (Ms = 5.4).

	Az	36	21	239	119	205	266	358	220	210	320	52	303	157	346	172	113	359	181	184	263	189	119	281	285	49	243	359	316	306
F	Dip	25	11 2	7 2	36 1	31 2	70 2	- س	48 2	24 2	3	42 5	65 3	43 1	52 3	57 1	65 1	52 3	27 1	22 1	55 2	27 1	57 1	49 2	73 2	67 2	64 2	43 3	48 3	52 3
-	Az D	24 2	286 1	339	19	332 3	89 7	89	34 2	339 2	100	285 4	115 6	7 2	117 5	343 5	304 6	141 5	291 2	290 2	168	43 2	28 5	53 4	181 7	208 6	14 6	219 4	47 2	200 5
œ	Dip A	23 2	24 28	54 33	14 1	45 3.	20 8	20 8	42 3	55 3;	79 10	34 28	25 1	43	27 1	33 34	24 3(32 14	34 29	35 29	4 16	58 4	1	30 5	4 18	21 20	17 1	39 2	4	13 2(
-	Az D	22 2	134 2	145 5	271 1	96 4	358 2	266 2	127 4	108 5	229 7	173 3	206 2	261 4	220 2	76 3	212 2	244 3	62 3	69 3	75 4	287 5	297	159 3	90 2	301 2	110 1	111 3	138	101 1
٩	Dip A			35 14		29 9	35	70 26	3 11	24 1(7 21	30 17	3 2(16 26	24 22	1.2		19 24	43 6		35 7	15 28	33 29	25 15	16 9	7 30	18 1	22 1	42 13	36 1(
		21	63		51						_	10000			2	4	4			d. 47	6		3	25.410	1		1	s. 2	4	3
_	e type	20	s n.d.	3 d.n.d.	n.S.	d.i.d.	i.S.	8 n.d.	i.s.	d.d.	d.i.d.	d.i.s.	i.d.	i.d.	i.S.	i.s.	i.d.	i.d.	n.s.	ð d.n.d.		d.i.s.	1	i.d.	1	i.s.	1	d.i.e	1000	122
2	Rake	19	-118	-158	-31	179	83	-118	40	180	179	12	125	135	62	50	121	149	-55	-156	92	10	87	154	95	64	71	20	103	103
PLANE	Dip	18	60	61	16	45	48	48	61	55	79	35	47	73	74	57	54	38	81	39	80	59	12	34	62	56	65	42	3	82
ш	Strike	17	271	289	318	240	69	69	4	249	ъ	210	320	204	109	301	144	14	118	230	348	331	25	293	3	193	7	155	240	22
i	type	16	n.s.	n.s.	n.d.	i.s.	i.d.	n.s.	i.d.	i.s.	d.i.s.	i.d.	i.s.	i.s.	i.d.	.d.	i.s.	i.s.	d.n.d.	n.s.	i.s.	.d.		i.s.	i.	i.d.	i.d.	i.d.		i.s.
	Rake	15	-51	-31	-63	45	117	-63	144	35	11	124	58	23	149	135	55	56	-164	-54	69	149	91	59	81	123	125	130	89	33
PLANE 1	Dip	14	40	71	82	89	49	49	56	90	89	83	53	48	32	50	46	71	36	75	11	82	78	76	29	42	31	77	87	15
2	Strike	13	138	187	17	331	287	287	252	339	95	110	94	309	352	196	278	129	221	121	147	236	208	45	173	54	225	50	47	144
	Mw S	12	5,4		е С	3,2	3,1	3,5	3,4	2,6				2,9	5	3,1		2,6	3,2	3,1	e e e e e e e e e e e e e e e e e e e	2,7		3		2,5	3,2	3,2		2,6
	ML	11					4	4,2									-													
	MD	10														4,1											_			
	Ms	თ		ъ	3,9								2,9			3,4	3		3											
	Uepth (km.)	œ	40	42	11	34	25	2	19	33	10	15	13	10	33	10	10	10	4	0		10	33	10	33	10	4	33	10	38
	Long.E	7	27,00	26,70	26,12	27,38	25,89	25,34	26,37	25,70	22,70	26,73	22,71	27,18	23,15	27,11	27,04	27,17	27,14	27,46	25,49	26,91	25,97	26,31	26,31	26,59	25,45	25,76	26,64	27,86
	Lat.N	و	44,60	44,90	44,18	44,07	44,55	44,75	43,98	44,23	44,79	43,80	44,40	44,74	44,33	44,47	44,52	44,34	44,54	44,78	44,64	44,69	44,47	44,43	44,58	44,60	44,19	43,99	43,58	44,66
Origin	ŝ	5	12:51:52	21:00:42	21: 16: 15	19:51:18	07:59:04	19:01:05	04:21:19	04:54:42	19:12:32	03:06:33	11:00:59	12:43:20	22:01:36	00:03:44	04:25:00	13:07:48	09:34:20	12:57:14	05:17:57	18:26:42	03:50:15	10:12:31	02:00:03	12: 19: 08	11:35:09	01:41:48	07:26:25	11:27.03
	Day	4	4	27	20	8	25	7	12	21	23	4	17	31	28	29	4	30	15	15	26	28	22	7	12	5	31	7	26	7
	Month	ε	1	2	4	12	8	7	9	80	00	10	2	5	6	б	10	ۍ	7	7	6	<i>с</i> о	6	11	5	8	12	2	3	10
-	Year	2	1960	1967	1977	1980	1983	1988	1990	1991	1991	1991	1992	1993	1993	1993	1993	1994	1994	1994	1995	1995	1996	1997	1999	2000	2000	2001	2002	2002
-	No.	F	1	2 1	3	4 1	5 1	6	7 1	8	9	10 1	11 1	12 1	13 1	14 1	15 1	16 1	17 1	18 1	19 1	20 1	21 1	22 1	23 1	24 2	25 25	26 2	27 2	28 2
<u> </u>																														

The focal mechanisms solutions of some earthquakes occurred in the Moesian Platform and Carpathian Foredeep

Table 2

(p				in the second		Terror 1	Sec. 1	and a		2004					452				(hans)						_						1254		
tinue	8	297	62	325	190	290	252	166	108	248	60	55	223	220	129	4	27	59	113	92	211	77	115	236	30	272	297	200	20	212	188	264	132
con (35	34	50	18	35	43	88	46	33	77	\$	8	21	43	28	27	2	61	75	9	37	14	47	12	59	41	36	47	39	65	30	38	õ
Table 2 (continued	24	173	245	124	69	91	17	266	239	35	245	277	60	10	354	141	199	242	338	278	76	258	14	73	206	40	96	95	214	2	354	110	б
T	23	40	40	71	37	45	7	10	45	11	36	62	68	43	54	55	22	29	10	8	43	76	11	22	31	35	52	13	50	22	59	49	Q
2	3	52	154	233	308	191	107	9	359	126	153	152	315	116	231	263	290	151	246	182	321	167	274	327	297	153	199	354	115	96	94	4	272
5	21	32	~	9	34	10	2	42	27	7	2	17	9	16	21	21	3	~	10	~	25	0	41	ю	2	30	11	40	7	11	7	13	~
	8	d.i.s.	i.s.	d.i.s.	i.d.	i.s.	, L	d.i.d.	d.i.d.		i.d.	d.i.s.	i.d.	i.s.	i.d.	i.s.	i.d.	i.s.		d.i.d.	i.d.	d.i.d.	-	d.i.s.	i.s.	i.S.	i.s.	d.i.s.	i.d.	i.d.	i.s.	d.i.s.	<u></u>
	19	2	40	17	127	41	88	167	175	107	134	4	160	44	144	35	121	51	103	175	133	170	100	~	51	55	35	19	145	128	27	20	82
	8	40	57	82	90	69	47	11	45	39	57	62	80	73	86	85	46	51	56	86	83	80	87	79	54	84	73	14	69	39	74	54	39
	17	85	210	100	249	69	16	163	141	227	93	195	268	353	179	135	42	215	165	136	263	121	194	12	181	35	73	24	61	213	325	51	356
	16	i.d.	i.d.	d.i.d.	d.i.s.	i.d.	81	1	i.s.	E E	i.s.	i.d.	d.i.s.	i.d.	d.i.s.	d.i.d.	i.S.	i.d.	8	d.i.s.	d.i.s.	d.i.s.	d.i.s.	d.i.d.	i.d.	d.i.d.	i.d.		i.s.	i.s.	d.i.d.	i.d.	
	15	130	139	172 0	1	153	92	80	45	76	43	152	11 0	157	5	175 0	61	128	72	4	11	10 0	15 0	169	131	169 (160	103	25	64	162 (142	97
	14	89	58 、	73 `	37	53 5	43	88	86	53	53	87	71	48	54	55 、	52	53	36	85	44	80	11	84	51 、	36 `	57	86	58	60	64	74 、	52
	_		_		-	· · · ·	05. 5					103	2	18			180								_						15		
	2 13	6 354	8 96	6 8	9 340	8 321	9 198	7 266	6 235	7 26	8 212	10	6 1	8 248	3 272	8 42	18	5 87	8 323	227	-	213	299	281	55	9 296	332	1 276	165	348	7 227	309	187
	11 12	2,	2,8	2,1	2,	2.	2	2.	2,	2,	2,8		2,1	2	2.	2,8		2.	2		_	-			_	2		3,	n		2,		_
	10											<u>б</u>								<u>ත</u>	2	2,2	6	2,7	2		2,8	_	2.7	8		8	2,9
	6		-									2							_	-	2		(1	(1	_		(1	_	2,	5		2	- (N
	ω	20	0	16	2	10	10	33	7	56	33	10	2	11	0	6	5	0	0	0	10	10	11	0	10	7	10	10	5	0	3	4	0
	7	26,30	26,71	26,52	27,18	26,49	26,52	26,59	26,55	25,98	25,84	27,01	26,85	25,44	27,74	25,56	24,84	27,47	26,88	28,27	27,00	28,41	26,87	26,77	27,03	26,98	27,03	27,01	27,91	27,02	27,00	26,96	27,03
	_													_						_	_							_					
	9	44,56	43,60	44,24	44,31		44,22	44,74	44,30	44,02	44,63	44,71	44,47	44,74	44,36	44,16	43,90	44,25	44,41	44,25	44,64	44,26	43,81	44,69	44,76	44,69	44,80	44,78	44,31	44,74	44,75	44,67	44,79
	5	09:14:25	20:34:20	08:22:44	06:39:12	11:34:50	00: 13:52	03;03:27	23:43:16	12:05:36	01:33:10	12:41:09	12:12:17	20:22:54	11:08:25	03: 14: 24	10:43:12	20:38:17	00:01:25	11:20:24	05:04:43	08:42:03	03:53:47	01:22:19	17:07:24	17:20:25	21:18:04	17:53:55	12:23:23	04:13:55	12:18:09	20:00:00	22:55:42
	4	20	16	3	5	7	10	27	7	27	2	4	10	21	4	18	5	12	12	15	9	17	8	22	24	24	24	25	26	27	27	27	27
	e	10	1	2	3	3	3	3	4	8	11	12	7	80	6	9	10	10	2	ю	4	4	5	00	00	00	8	00	8	00	8	ω	ω
	2	2002	2003	2003	2003	2003	2003	2003	2003	2003	2003	2003	2004	2004	2004	2004	2004	2004	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005
	-	29 21	30 21	31 21	32 21		34 21	35 21	36 21	37 21	38 21	39 21	40 21	41 2	42 21	43 21	44 21	45 21	46 21	47 21	48 21	49 21	50 21	51 2	52 21	53 21	54 21	55 21	56 21	57 21	58 21	59 21	60 21
		3531	335	689)	123	3238	2952	2/2	- 2006	73535	253	313	8	37	2250	- 20	37	826	22	1	8	1	80	0225	102	2933	0220	- 2022	2532	12	1232	5755	- 23

8	86	283	148	306	289	39	254	250	87	211	186	204	88	285	352	normal -lateral
25	57	53	37	36	37	53	2	9	33	80	25	67	8	49	47	- left
24	294	17	358	181	91	231	78	56	311	315	333	331	317	129	90	ult; n .n.s
33	32	4	49	39	52	36	26	84	48	2	61	14	14	38	7	eral fa lip; d
3	199	110	250	62	192	137	347	160	193	45	6	99	223	29	187	ht-late cike-sl
3	2	37	15	31	0	9	2	~	23	10	14	18	16	12	42	al rig ral sti
8	i.d.	Ē.	i.d.	i.d.	i.s.	i.d.	i.d.	d.i.d.	d.i.s.		i.s.	i.d.	20	i.d.		- norm nt-late
5	128	87	140	129	33	133	124	177	6	8	38	119	106	133	82	n.d. – – rigł
18	60	82	76	87	72	60	52	85	48	35	83	30	62	67	87	e fault; d.n.d.
17	136	17	195	2	66	76	280	294	234	138	319	176	145	328	90	– reverse right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; n.d. – normal right-lateral fault; n.s. – normal right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; d.n.d. – right-lateral strike-slip; d.n.s. – left-lateral in A i.s. – left-lateral strike-slip.
16	i.s.	i.d.	d.i.s.	d.i.s.	i.d.	i.s.	i.s.	d.i.s.	i.d.	1	d.i.d.	-	i.s.	i.s.	i.d.	t-lateral obliqu
15	43	113	18	5	158	40	55	5	138	87	172	74	62	32	158	erse lef -lateral
14	47	6	52	39	58	50	49	87	84	55	62	64	32	47	8	s. – revo erse left
13	258	220	296	26	324	194	53	25	138	313	226	324	293	81	338	ault; i.s – reve
12						2,9								2,6		ique f lt; i.s.
11															2,9	e fau
9	3,1	3	3,1	2,8	2,7		3,1	2,7	3,1	2,7	2,4	2,6	3,6		-	t-later obliqu
თ												_				teral
8	2	0	10	10	80	6	9	1	9	0	15	0	30	89	15	revers ight-la
7	27,06	27,06	26,99	27,23	26,82	25,85	26,70	28,17	26,58	28,00	28,27	27,16	28,08	26,59	26,27	lt; i.d. – everse r eite clin
9	44,76	44,78	44,69	44,78	44,75	44,62	44,29	44,56	44,12	44,68	44,02	44,64	44,73	43,98	44,73	erse faul i.d. – r
5	23:50:52	06:17:06	15:59:07	09:46:18	16:33:26	14:50:40	11:14:19	15:30:05	16:54:50	16:09:46	03:43:13	12:14:03	10:47:06	00:45:11	03:58:24	Legend: n. – normal fault; i. – reverse fault; i.d. – reverse right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; n.d. – normal right-lateral fault; n.s. – normal left-lateral oblique fault; i.d. – reverse right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; d.n.d. – right-lateral strike-slip; d.n.s. – left-lateral strike-slim: d.d. – right-lateral strike-slip; d.e. – left-lateral strike-slip; d.n.s. – left-lateral
4	27	28	28	3	15	27	16	31	31	10	14	5	9	17	24	al faul I oblig
ε	ω	8	8	6	6	9	10	10	10	11	11	12	2	6	4	– norm ft-latera ite-clin
2	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2005	2006	2006	2007	end: n. lef
-	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	Leg



Fig. 6 – The dip (a) and azimuth (b) of P and T axes of earthquakes occurred in the Moesian Platform.

4. NORTH-DOBROGEAN OROGEN

The seismicity of this region has been observed since 1949 by Ion Atanasiu, who evidenced the lines of "seismic sensitivity": Tulcea–Isaccea– Galați, Cerna–Măcin–Brăila and Babadag– Nicolae Bălcescu. In a previous paper (1938), Ioan G. Popescu mentions the epicenters from Tulcea, Isaccea, Babadag and Topolog. There were located, between 1871–1929, 11 earthquakes in the Galați–Sulina area and 3 shocks connected by the Peceneaga–Camena fault.

The information presented by us refers to the area included within the following geographic coordinates: latitude 44.5° – 45.5° N and longitude 28.0° – 29.0° E. In the south, the data included a part of Central Dobrogea too, belonging to the Moesian Platform.

Thus, between 1831-2006, were located 218 earthquakes with M = 1.9 to 5.6 and depths of 0–33 km, in the region; 51 shocks had magnitudes greater than 3.5 (Figs. 7, 8).

Historical earthquakes produced in the XIXth century and early XXth century had magnitudes and depths estimated at higher values; so, seismic events produced at depths of 40–50 km and a magnitudes of 5.3 to 5.6 existed. Epicenter distribution indicates their concentration in SE Măcin, Tulcea area north of Babadag.

In Table 3 are presented the focal mechanisms of 18 earthquakes produced during 1980–2006. The faulting plane indicates two average directions: one, N35⁰E (NE–SW) and another, N43⁰W (NW–SE), with a predominant dip of 56⁰. The faulting type was the reverse fault (16) and two strike-slip faults. The average tectonic stress directions had been as follows: for

compressive forces (P), N63⁰E (ENE–WSW) and N44⁰W (NW–SE), the latter being predominant, the tension forces (T), a main direction, N41⁰W (NW–SE) and other, secondary, N60⁰E (ENE–WSW). The compressive forces operated in a horizontal plane (with an average slope of 18°) and the tension forces had an average dip of 60° (Fig. 9).

The strong earthquake occurred on November 13, 1981 (Mw = 5.1), located east of Tulcea, was generated on a fault oriented NE–SW, situated at depth of 4–9 km (Oncescu *et al.*, 1989). Main shock (MD = 5.4) was followed by six aftershocks with MD = 2.9 to 3.5. The authors noted the migration of seismic activity from NE to SW along a distance of about 23 km.



Fig. 7 – Earthquake epicenters distribution that occurred in Dobrogea during the 1831–2006 period (tectonic background after Visarion *et al.*, 1990).

The recalculating of the faulting plane indicates a plane oriented NE–SW (N120⁰W). A similar solution provided Gabriela Polonic in 1986. The P axis was oriented ENE–WSW (N73⁰E) and the tension axis (T) NNW–SSE (N11⁰W).

The Niculițel–Cataloi–Cerna area revealed as a active seismic area with foci situated in the crystalline crust seismic of the Tulcea tectonic zone and of the Niculițel nappe. Many earthquakes are related to tectonic line Luncavița–Consul; foci were located at depths of 5–10 km, the area belonging to the crystalline crust.

The earthquakes located in the south of Tulcea are extended to the west, near the tectonic contact with the Măcin unit (line Luncaviţa–Consul). Some earthquakes are related to this tectonic contact or associated satellite faults. The region is particularly active there were located about 15 seismic shocks, concentrated in a relatively small area.



Fig. 8 – Earthquake epicenters distribution, isobaths map and 3-D image of the seismic foci of Dobrogea.

				Origin			4700					9	PLANE1		t)	P	PLANE 2	2	*!		٩	~	N		
Ś	Year	Month	Day	hh:mm:ss	Lat.N	Long.E	(km.)	đ	Mw	QW	ML	Strike	Dip	Rake	type	Strike	Dip	Rake	type	Dip	Az	Dip	Az	Dip	Az
-	2	ر	4	2	9	7	8	ი	6	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26
~	1980	0	11	23.24.25	45,32	28,03	20,4		4,2			338	0	17	d.i.s.	52	87	66		42	313	6	51	47	150
~	1981	11	13	09:07:13	45,17	29	15		5,1			22	56	66	i.s.	22	41	121	i.d.	~	141	20	48	69	252
ы	1991	ю	σ	15:21:36	45,04	28,81	10					191	8	136	, d	288	47	10	di.s.	24	247	46	4	35	139
4	1992	ô	3	20:21:43	45,2	28,9	10		3,1		3,9	237	83	100		35	29	71		17	320	9	52	70	168
ŝ	1994	11	24	10:24:39	45,15	28,41	10			3,3		274	37	175	d.i.d.	6	87	53	i.s.	32	128	37	10	37	246
9	1997	9	13	09:23:59	45,02	28,89	ъ					294	48	131	j.d.	62	56	54	i.s.	4	176	29	28	60	274
7	1997	00	21	08:49:42	45,12	28,36	10			3,1		55	22	102		171	13	27	i.s.	38	134	12	234	50	338
00	1998	8	ô	11:03:45	45,12	28,39	5			3,4		300	36	15	dis.	198	81	125	i.d.	28	260	35	12	43	141
6	1999	5	~	11:06:18	45,12	28,39	5			3,4		342	88	105		128	28	59	i.s.	20	61	14	156	666	278
10	2001	-	19	15:12:47	45,12	28,03	33					306	82	38	ai T	210	53	170	di.d.	19	72	51	316	32	175
11	2004	10	3	09:02:07	45,2	28,97	31	4,8				202	42	83	-	31	48	96		3	117	5	207	84	353
12	2004	11	~	08:55:55	45,22	28,24	0				3,1	73	65	62	5	305	37	135	i.d.	15	183	25	86	60	302
13	2004	11	24	18:36:55	45,16	28,79	3				3,8	273	18	160	i.d.	22	84	73		37	127	17	24	48	274
4	2005	-	18	17:15:30	45,14	28,88	0				3,8	252	55	149	i.d.	1	65	39	i.s.	9	124	45	28	45	221
15	2005	-	24	11:34:54	45,03	28,94	0				2,7	221	79	17	d.i.s.	128	74	168	di.d.	4	354	70	253	20	85
16	2005	4	18	12:29:21	45,18	28,94	5				3,5	223	89	172	d.i.d.	313	82	+	di.s.	5	268	82	36	6	178
17	2005	4	23	05:18:34	45,22	28,24	0				3,3	22	52	91		201	38	88	-	7	111	1	201	8	298
18	2006	2	9	10:47:06	44,73	28,08	30			3,6		293	32	62	S.	145	62	106		16	223	14	317	69	8
	n - Joneso I		. 1 1 U		1 J	L 2 .41.						1			-	E					-	-	-		

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Table 3

Legend: n. – normal fault; i. – reverse fault; i.d. – reverse right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; n.d. – normal right-lateral fault; n.s. – normal left-lateral oblique fault; i.d. – reverse right-lateral oblique fault; i.s. – reverse left-lateral oblique fault; d.n.d. – right-lateral strike-slip; d.n.s. – left-lateral strike-slip; d.n.s. – left-lateral strike-slip; d.n.s. – left-lateral strike-slip; d.n.s. – reverse left-lateral oblique fault; i.s. – reverse left-lateral oblique fault; d.n.d. – right-lateral strike-slip; d.n.s. – left-lateral strike-slip; d.n.s. – left-lateral strike-slip; d.n.s. – left-lateral strike-slip.

(a)

36





0

30

150

60

ò 330 330 3 -2 300 60 300 0 0 240 210 150 3 210 180 Р

(b)

Fig. 9 – The dip (a) and azimuth (b) of P and T axes of earthquakes occurred in North Dobrogea.

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